

A large, dramatic cumulonimbus cloud formation, also known as a mother cloud, dominates the sky. The cloud is massive and billowing, with a bright, glowing core in the center, suggesting intense upward air movement and potential precipitation. The sky is a deep, dark blue, and the city skyline in the background is silhouetted against the horizon. The overall scene is atmospheric and powerful.

Weather

🌐 Phases of water

- Latent heat
- Humidity

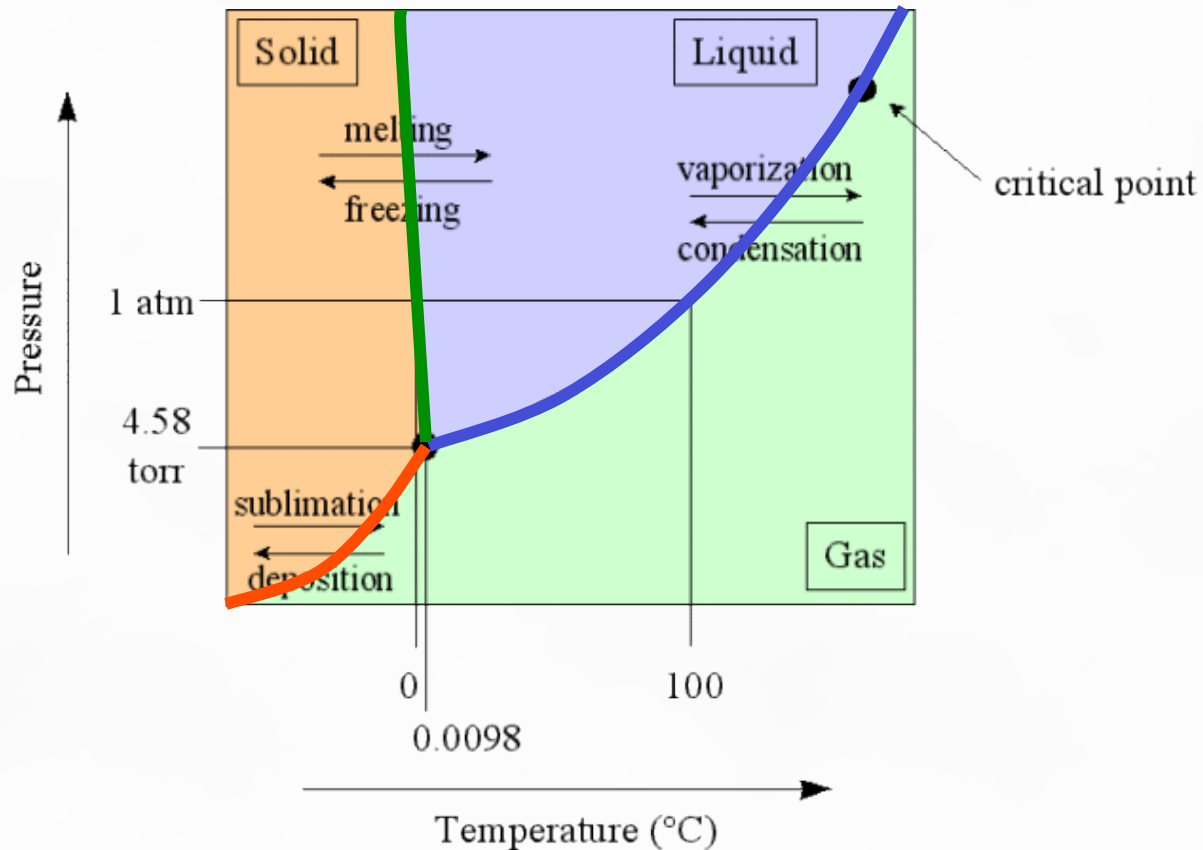
🌐 Cloud formation

- Adiabatic cooling

🌐 Air masses

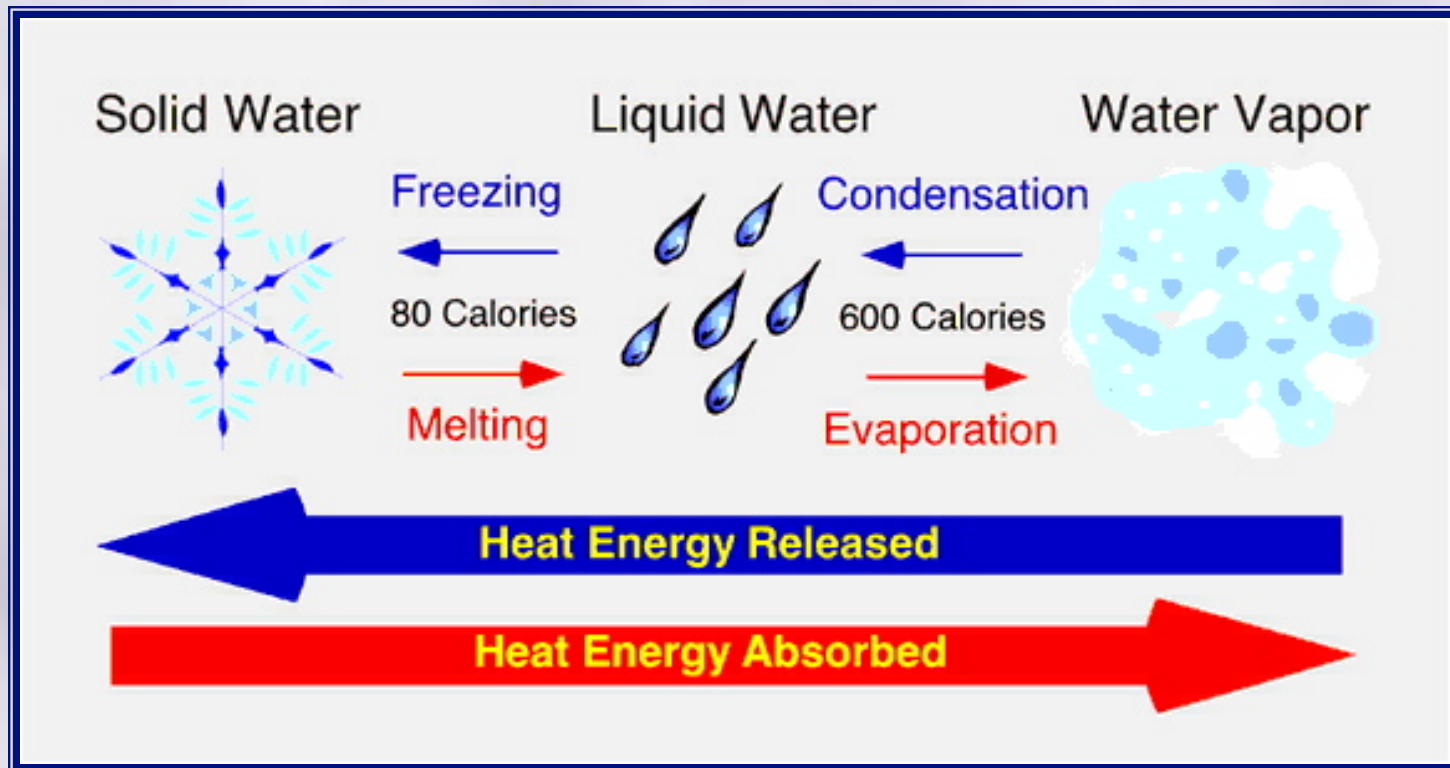
- Pressure
- Fronts

Phases of Water

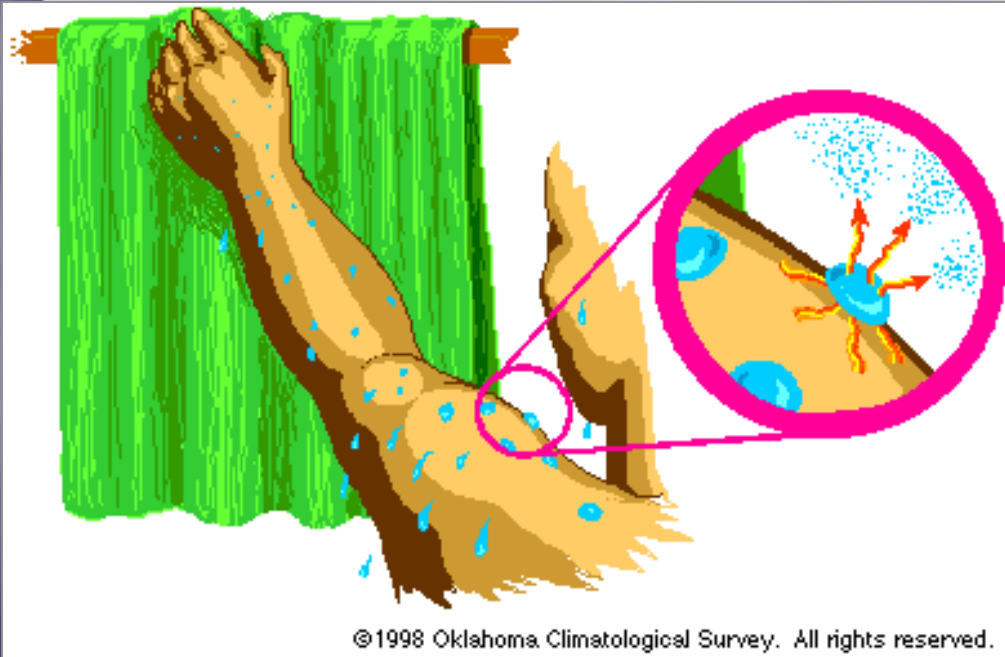


The phase (solid, liquid, or gas) of any matter depends on both temperature and pressure. Converting from one phase to another requires the absorption or release of energy.

Latent Heat



Latent heat – energy consumed or released during a *phase change*.



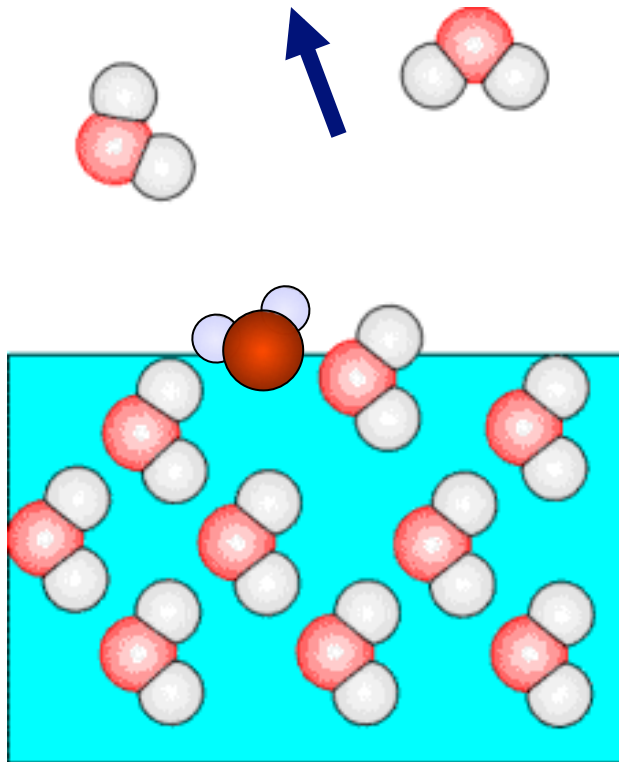
Latent Energy

Body heat is absorbed to change water from liquid to vapor.



As ice cubes melt, they absorb heat from the drink, cooling it.

Evaporation – water in liquid state escapes to become gaseous.



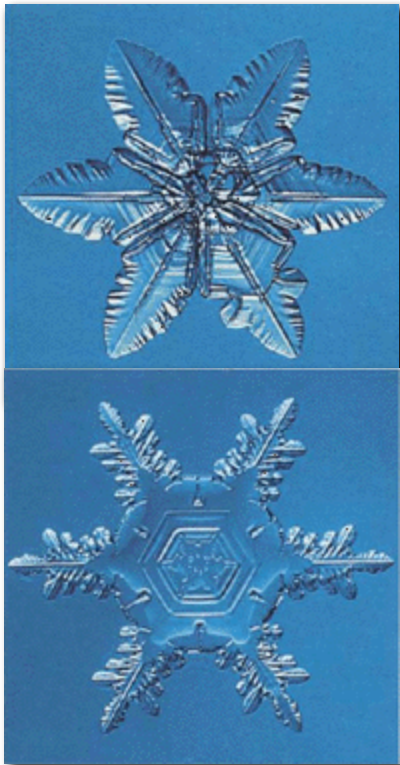
Latent heat is **absorbed** by evaporating molecules

Sublimation – molecules escape from solid to gaseous phase



Latent heat energy is **absorbed** by escaping molecules

Deposition – In the atmosphere, freezing vapor (gas) can form the beautiful crystals called snowflakes.



*Latent heat energy is **released** by depositing molecules*

<http://www.its.caltech.edu/~atomic/snowcrystals/>

Freezing – water in liquid state crystallizes to form a solid.



*Latent heat energy is **released** by freezing molecules*

<http://www.brockportfire.org/>

Condensation – water molecules in gaseous state stick together, forming liquid

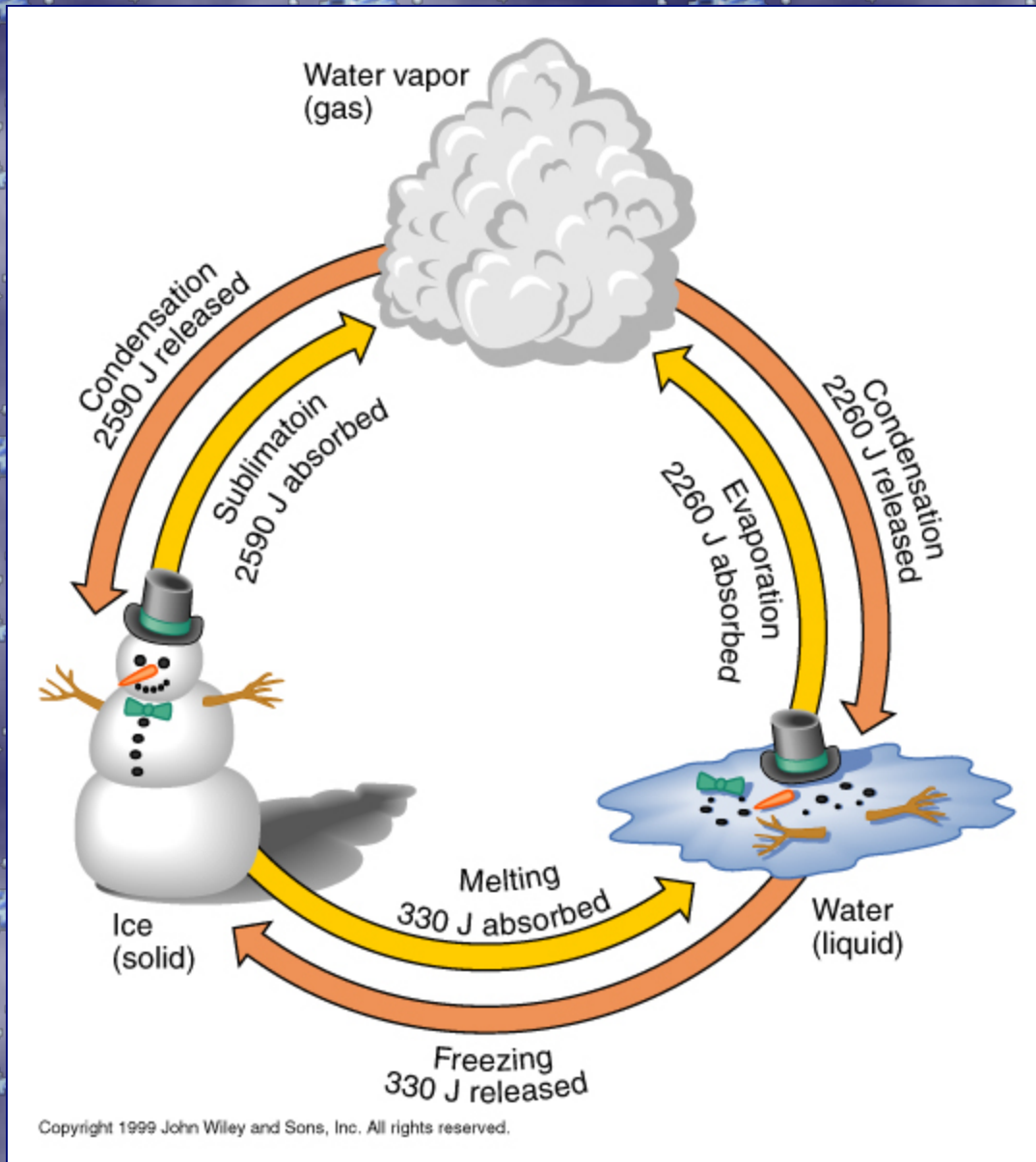


*Latent heat is **released** by condensing molecules*

In the atmosphere, condensation produces clouds and fog.



Condensation occurs more readily when the air contains aerosols, which act as condensation surfaces.



Humidity



Humidity – quantity of water vapor in a packet of air.

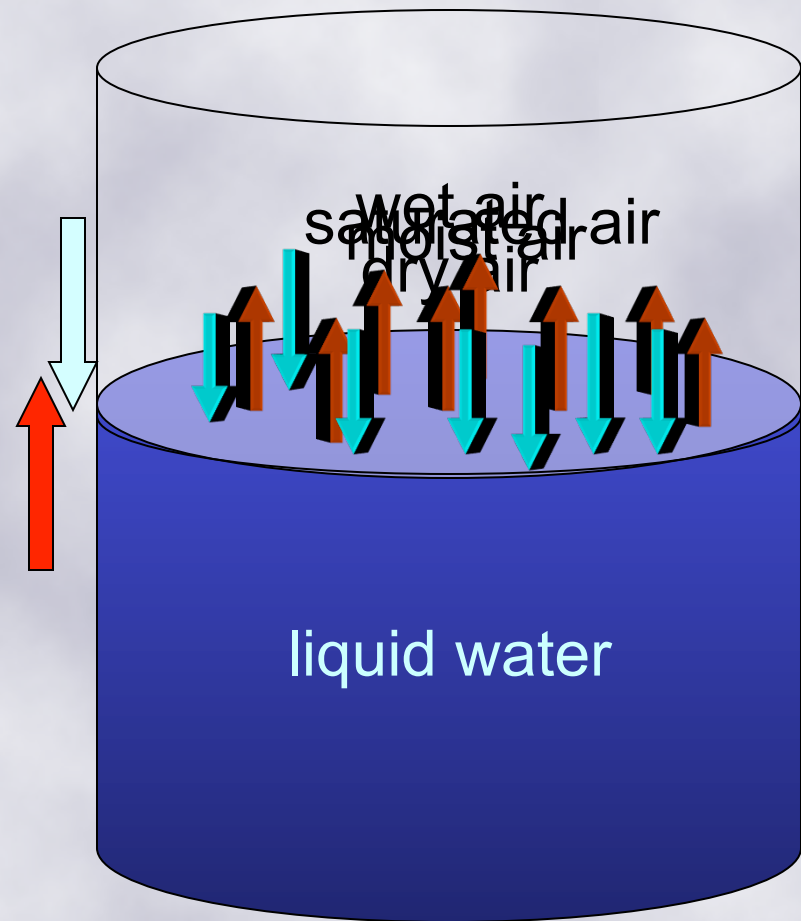
The actual amount of water vapor in the atmosphere rarely exceeds 4%. However, this “absolute” humidity is not a very useful measurement.

Instead, meteorologists measure *relative humidity*, which is the ratio of the measured humidity to the humidity at which condensation would occur for an air packet.

Humidity

Saturation – Point of dynamic equilibrium between water in liquid and gaseous form.

1. Water evaporates into the dry air, gradually increasing the water pressure in the formerly dry air.
2. As water in the air builds up, some makes the return trip, becoming liquid again.
3. When the amount of water evaporation equals the amount of air condensing, saturation is reached.



*Saturation is controlled primarily by **temperature***

Temperature	Saturation at:
-40°C (-40°F)	0.1 g _{water} /kg _{air}
0°C (32°F)	3.5 g _{water} /kg _{air}
20°C (68°F)	14.0 g _{water} /kg _{air}
35°C (95°F)	35.0 g _{water} /kg _{air}

A human body will feel much more atmospheric moisture at saturation at 95° than -40°.

Relative Humidity

Ratio of a parcel of air's measured water vapor and the saturation water vapor for that temperature

$$\text{Relative Humidity} = 100 \times \frac{\text{water vapor}_{\text{measured}}}{\text{water vapor}_{\text{at saturation}}}$$

Example: Parcel of air at 10°C (50°F), measured water vapor = 5 g_{water vapor}/kg_{air}

Water vapor at saturation (10°C) = 7 g_{water vapor}/kg_{air}

$$\text{Relative Humidity} = 100 \times \frac{5 \text{ g}_{\text{water vapor}}/\text{kg}_{\text{air}}}{7 \text{ g}_{\text{water vapor}}/\text{kg}_{\text{air}}} = 71.4\%$$

Dew Point

Temperature to which a parcel of air would need to be cooled to reach saturation

Example: Parcel of air at 10°C (50°F), measured water vapor = 5 g_{water vapor}/kg_{air}

Temperature at which 5 g_{water vapor}/kg_{air} is the saturation point (100% humidity) = 5°C (41°F)

Dew Point = 5°C

Cooling a parcel of air below its dew point results in condensation (water vapor → liquid water)

Cloud Formation

Clouds form when water vapor condenses to form liquid water droplets in the atmosphere

Condensation

Cool parcel of air below dew point

Adiabatic cooling

Adiabatic Cooling

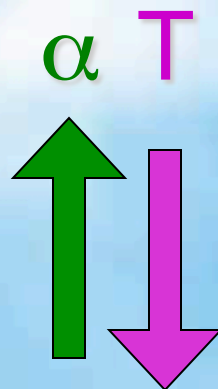
Drop in temperature brought about by change in volume

First Law of Thermodynamics

$$\Delta H = p(\Delta\alpha) + c_v(\Delta T)$$

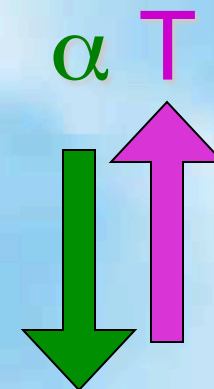
change in heat energy change in volume change in temperature

for adiabatic cooling, $\Delta H = 0$



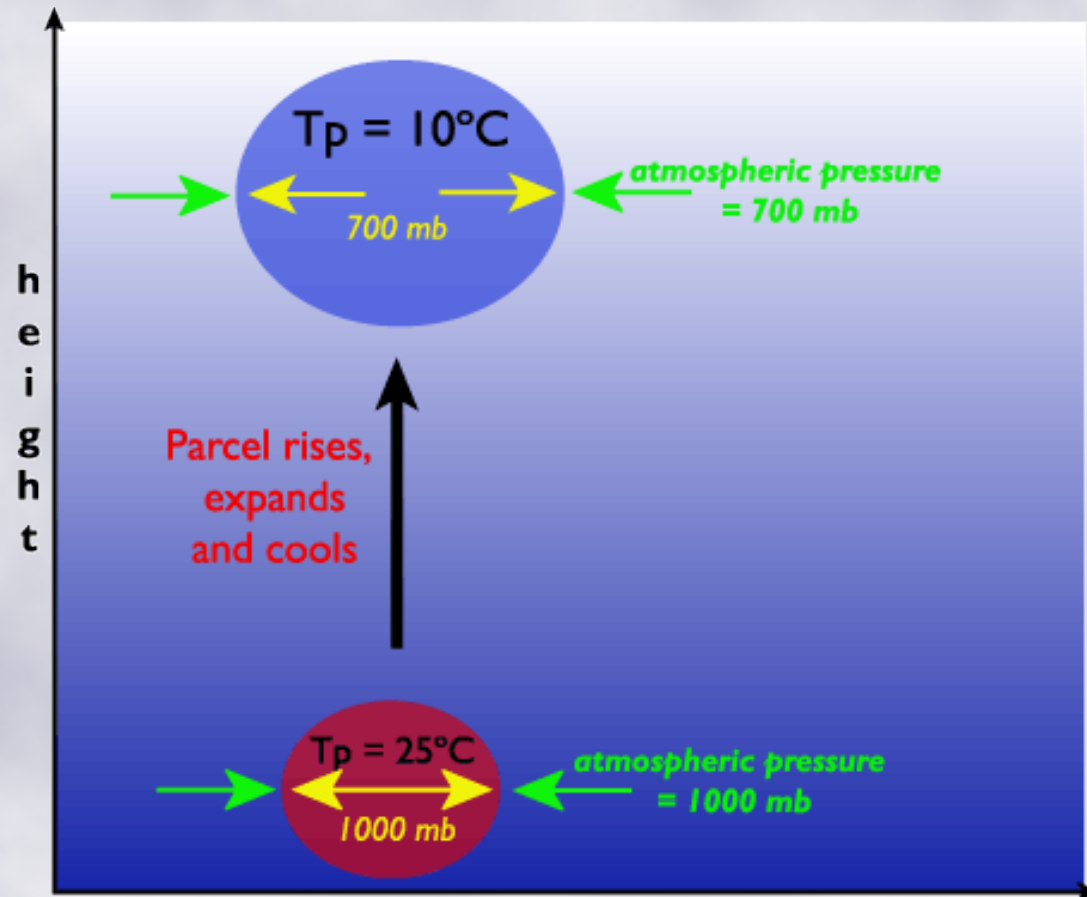
$$-c_v(\Delta T) = p(\Delta\alpha)$$

temperature volume



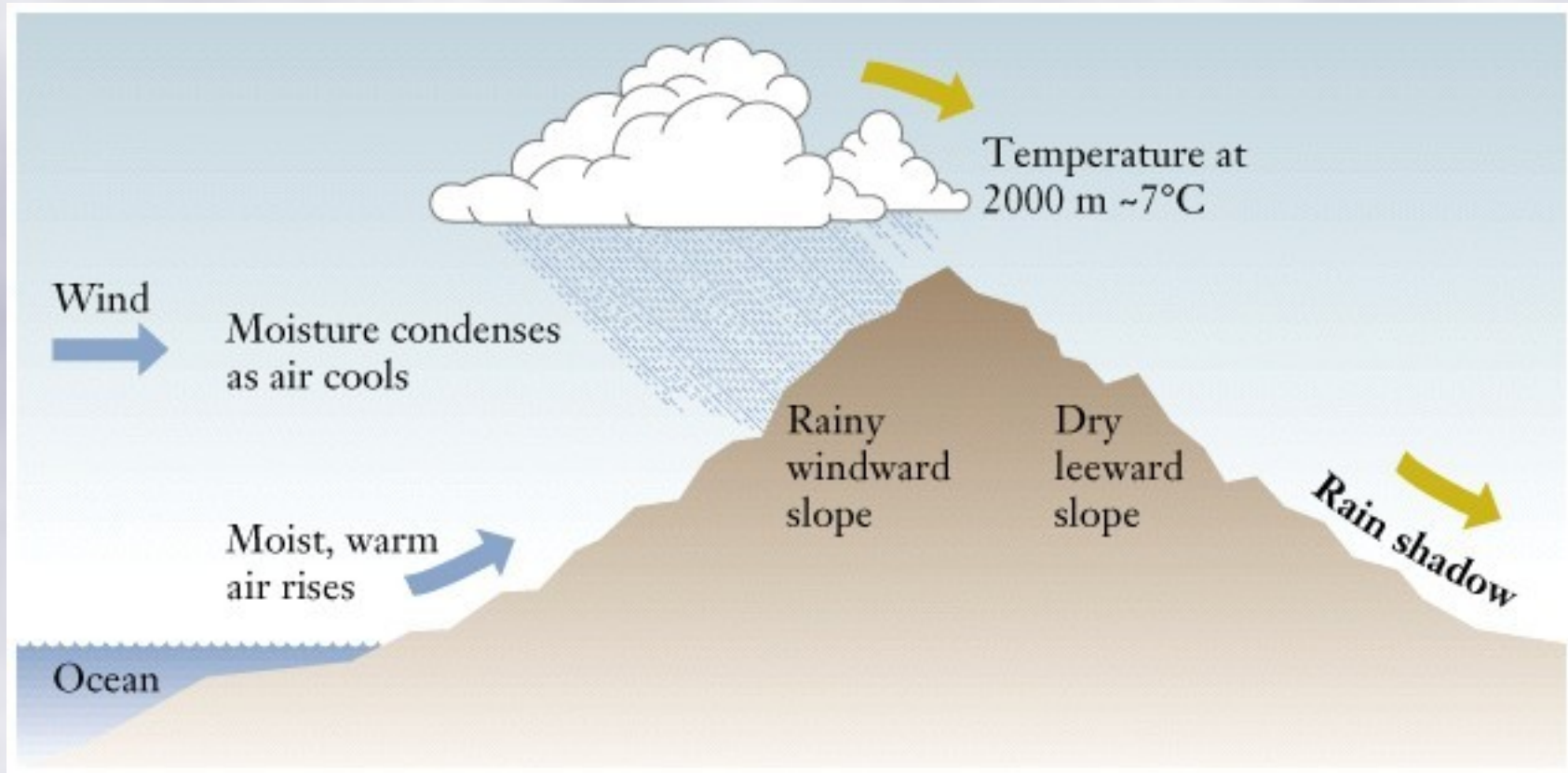
Adiabatic Cooling

As a parcel of air rises and expands, it cools



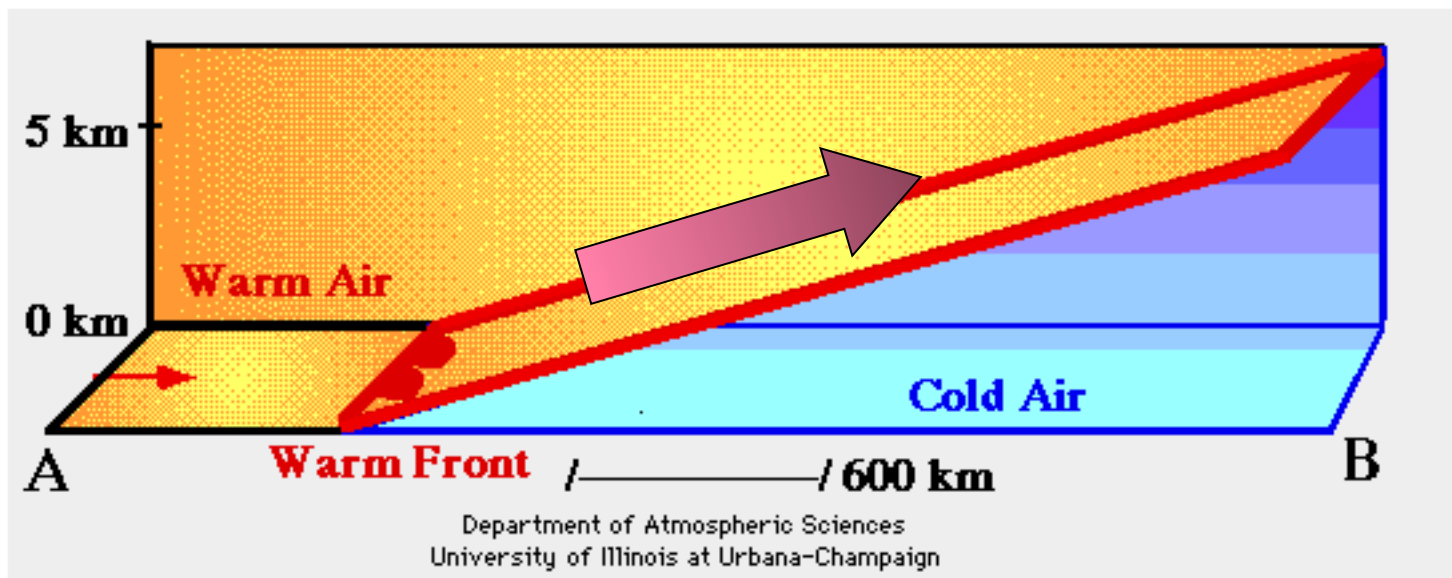
When the parcel has cooled to the dew point, condensation occurs and clouds form

Orographic Lifting

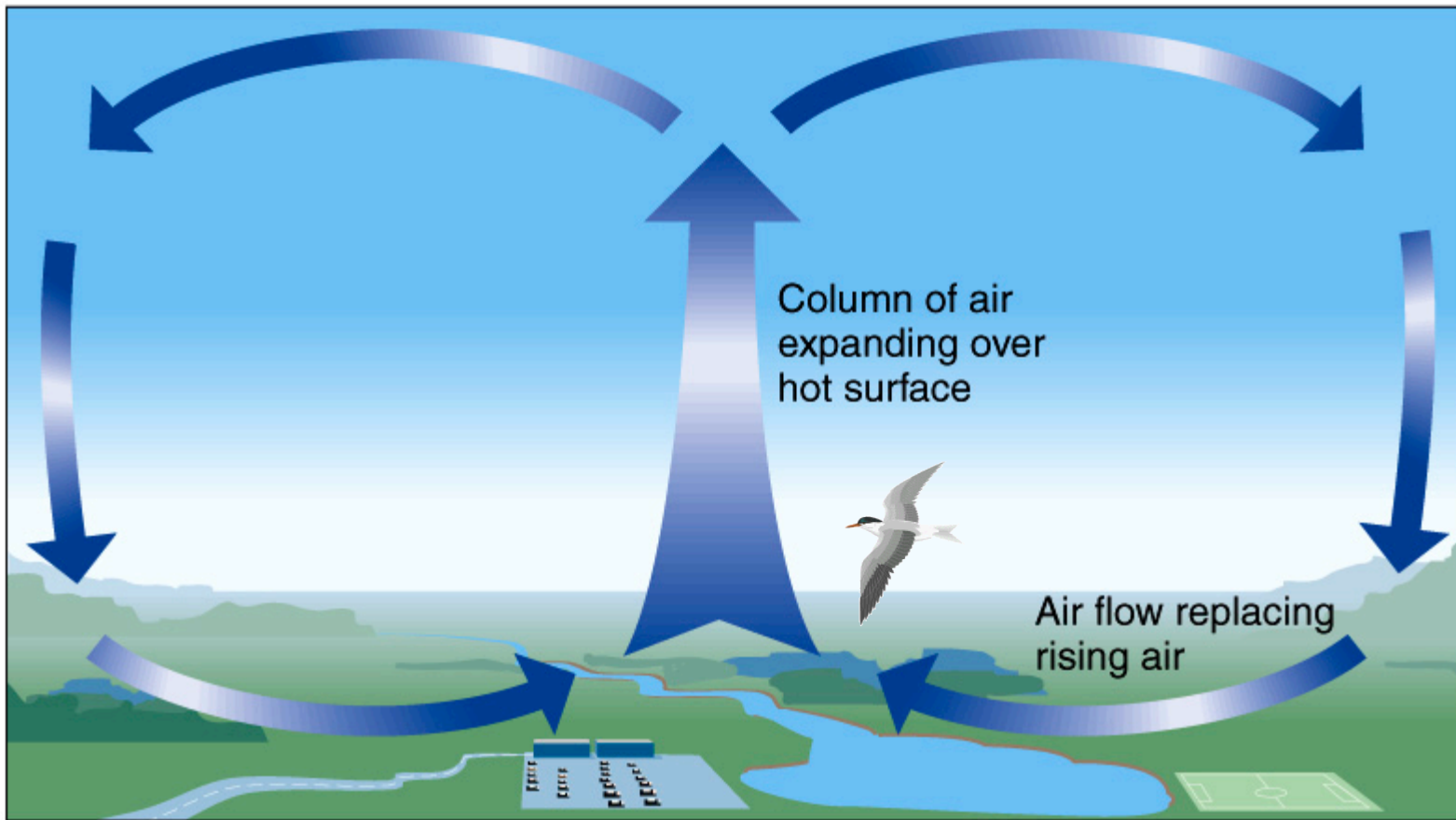


Frontal Wedging

warm parcel of air rises over cold parcel of air



Local Convection



Watch for birds spiraling upwards on “thermals” created by local convection off large parking lots

Atmospheric Stability and Instability

Parcels of air will rise as long as the temperature of the parcel is greater than the surrounding air.

For **stable air**, there is some elevation during rising where the internal and external temperatures equalize, and the parcel will not rise any further

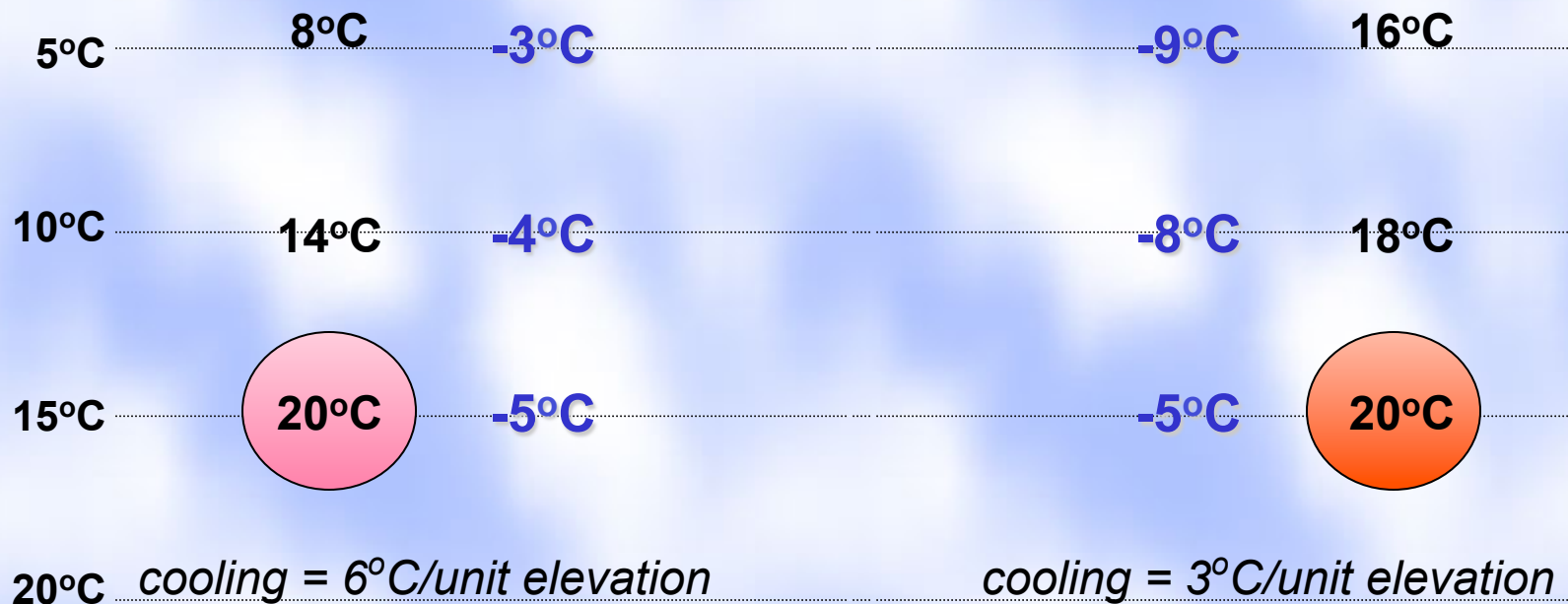
For **unstable air**, that equilibrium point is never reached by the rising air, and the hot parcel of air displaces the cold air above, causing turnover.

Stable and Unstable Air

As warm air rises, it expands and cools adiabatically

stable air

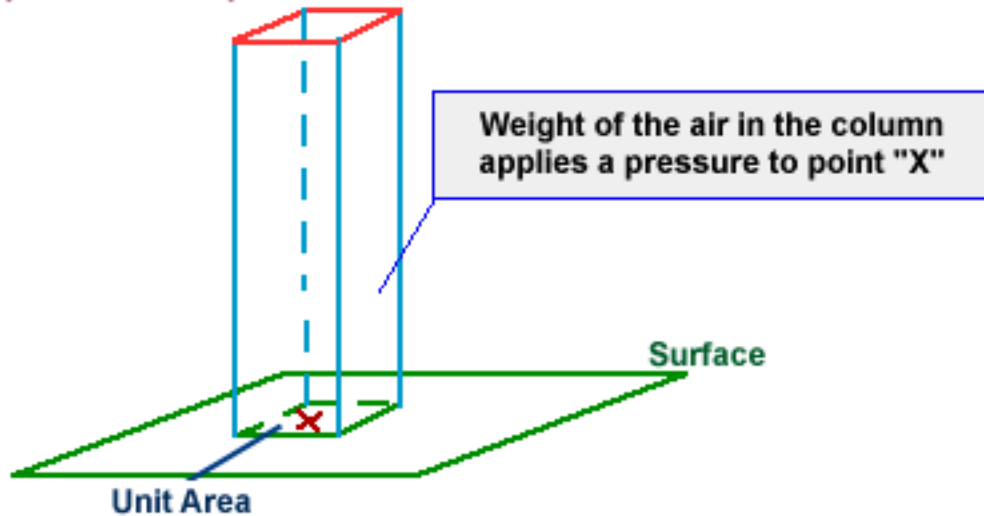
unstable air



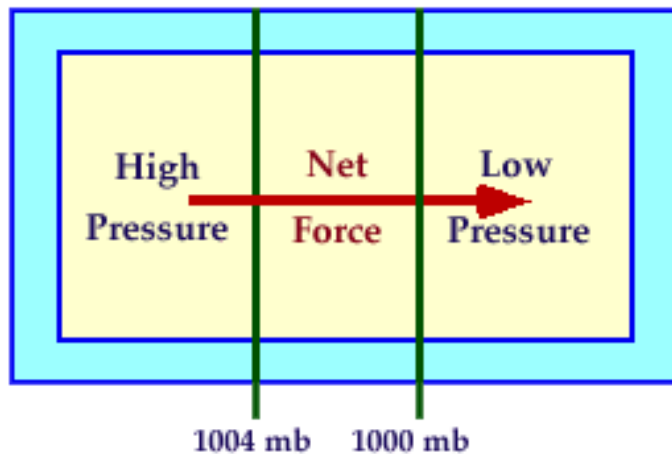
If the cooling is slower than the rate at which the surrounding air cools with elevation (the *environmental lapse rate*), then the parcel will remain warmer than the air through which it rises.

Atmospheric Pressure

Top of the Atmosphere



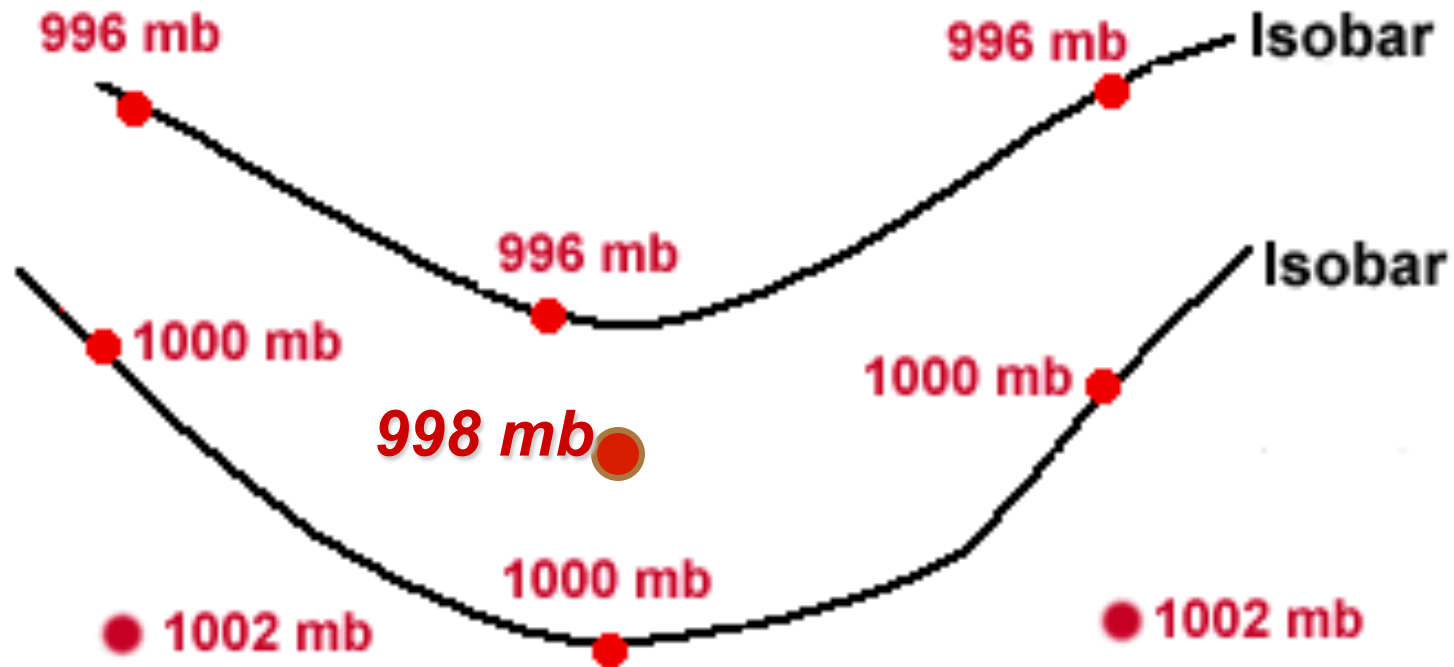
The amount of air “pressing down” determines pressure.



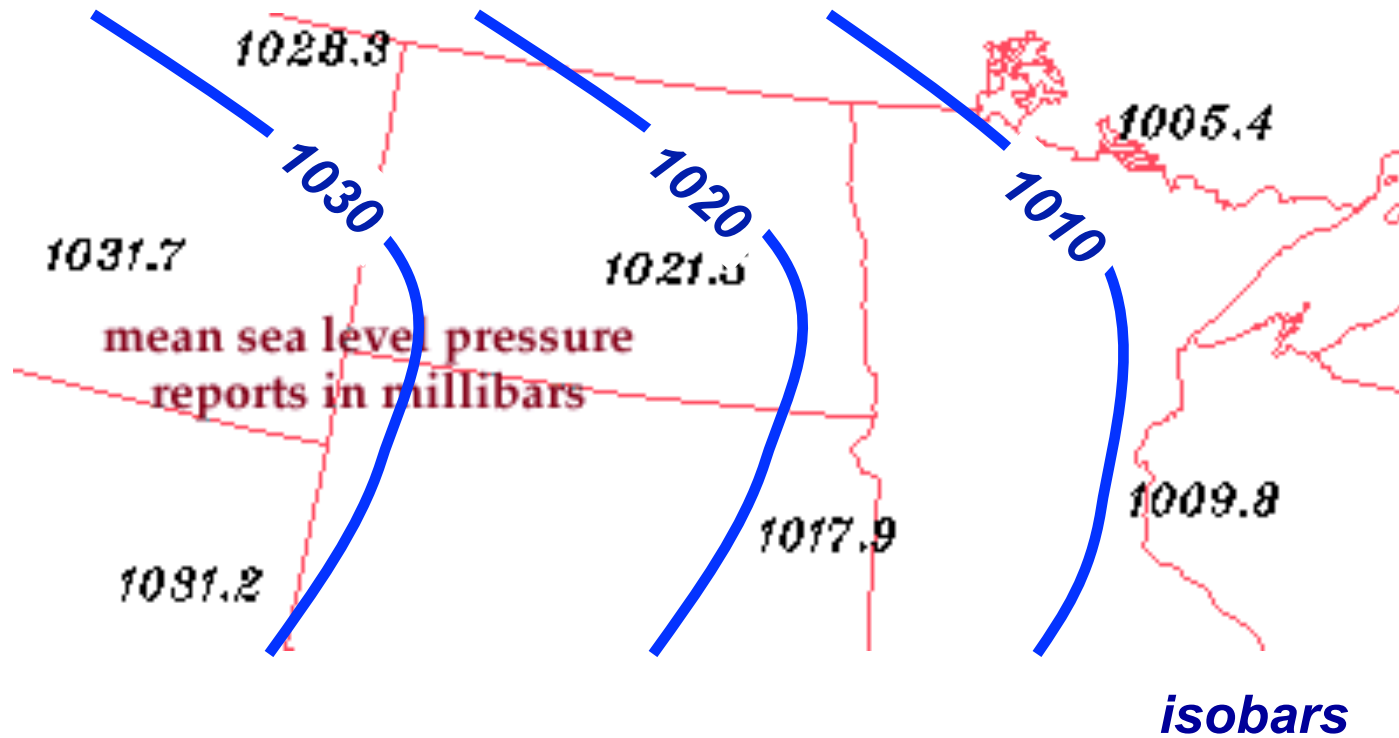
In general, air (and thus wind) flows from high pressure to low pressure.

Atmospheric Pressure

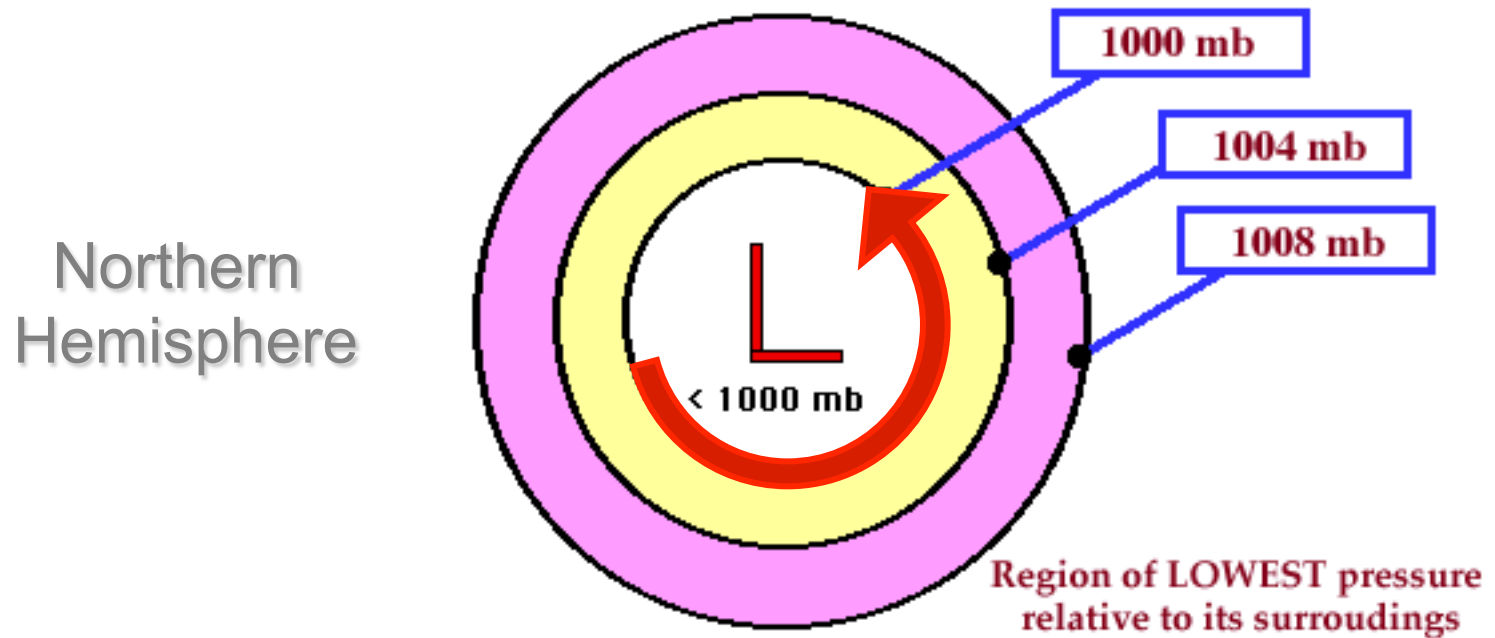
Weather maps usually display *isobars* (lines connecting points of equal pressure, like contour lines on a topographic map).



Atmospheric Pressure



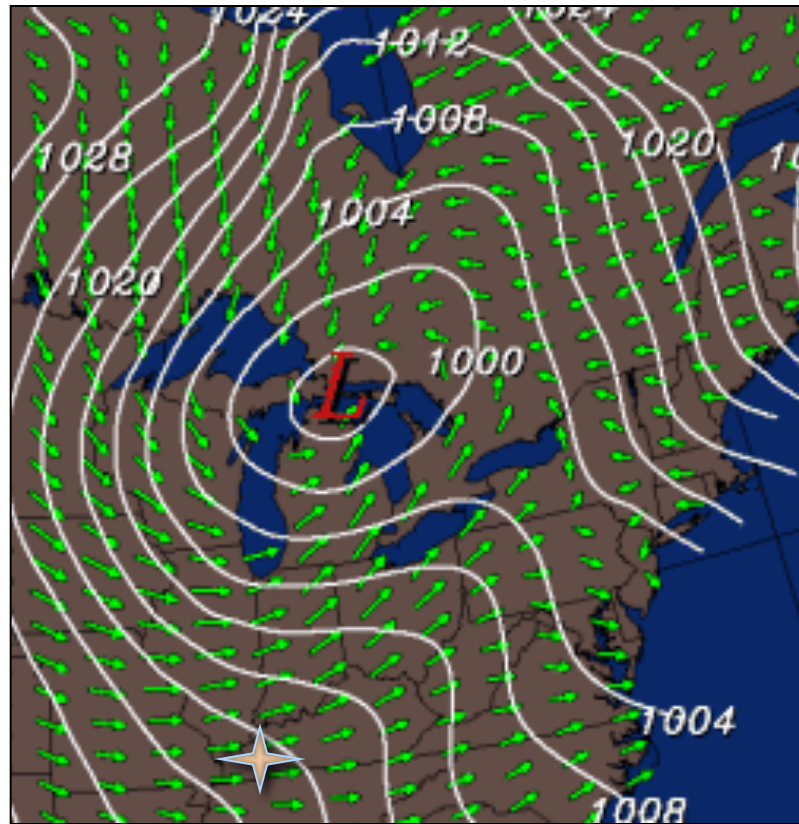
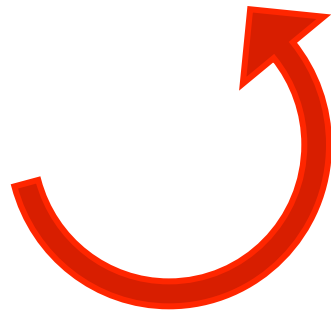
Low Pressure Systems



Winds tend to spiral in counter-clockwise to the center of a low pressure system (cyclone)

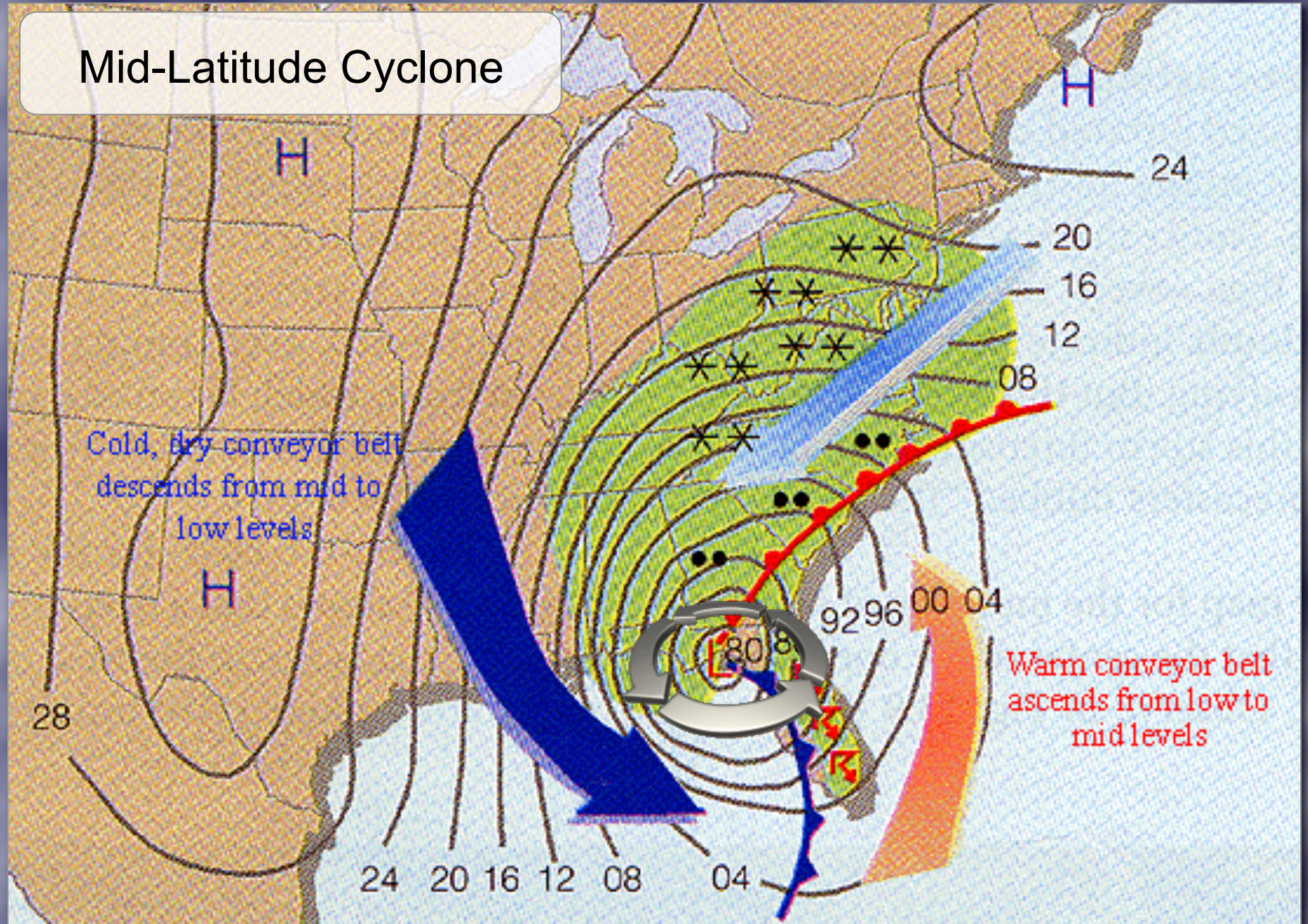
Low Pressure System

Northern Hemisphere

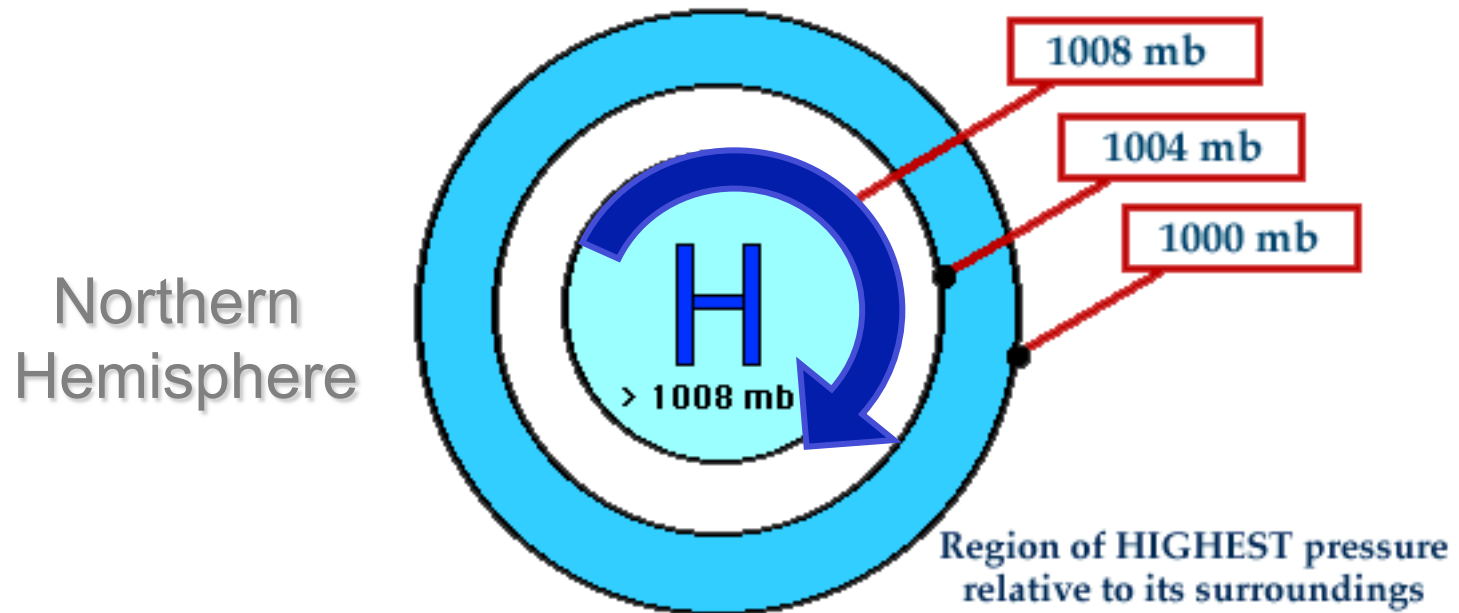


Winds tend to spiral in counter-clockwise from the center of a low pressure system (cyclone)

Mid-Latitude Cyclone



High Pressure System



Winds tend to spiral out clockwise from the center of a high pressure system (anti-cyclone)

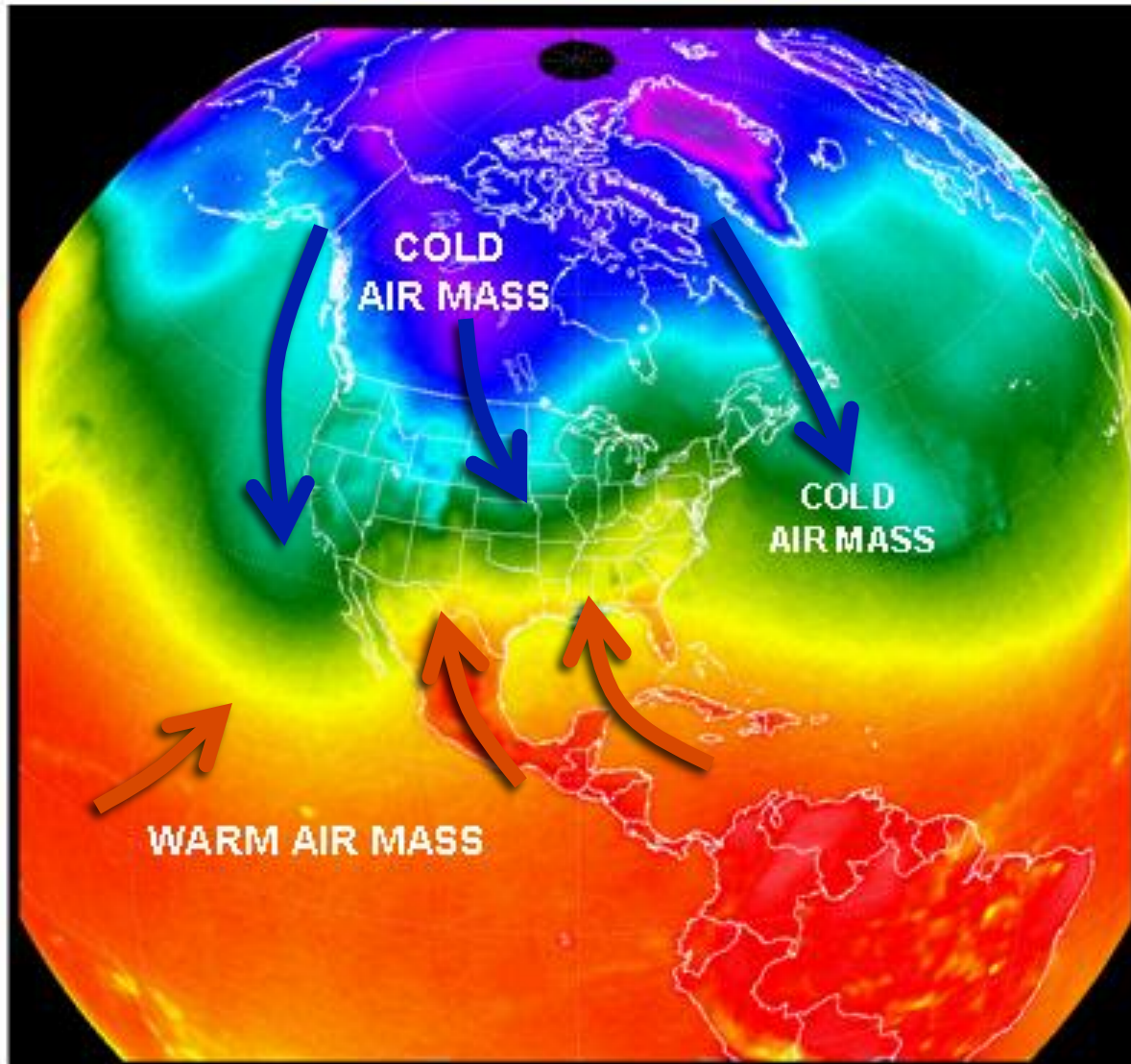


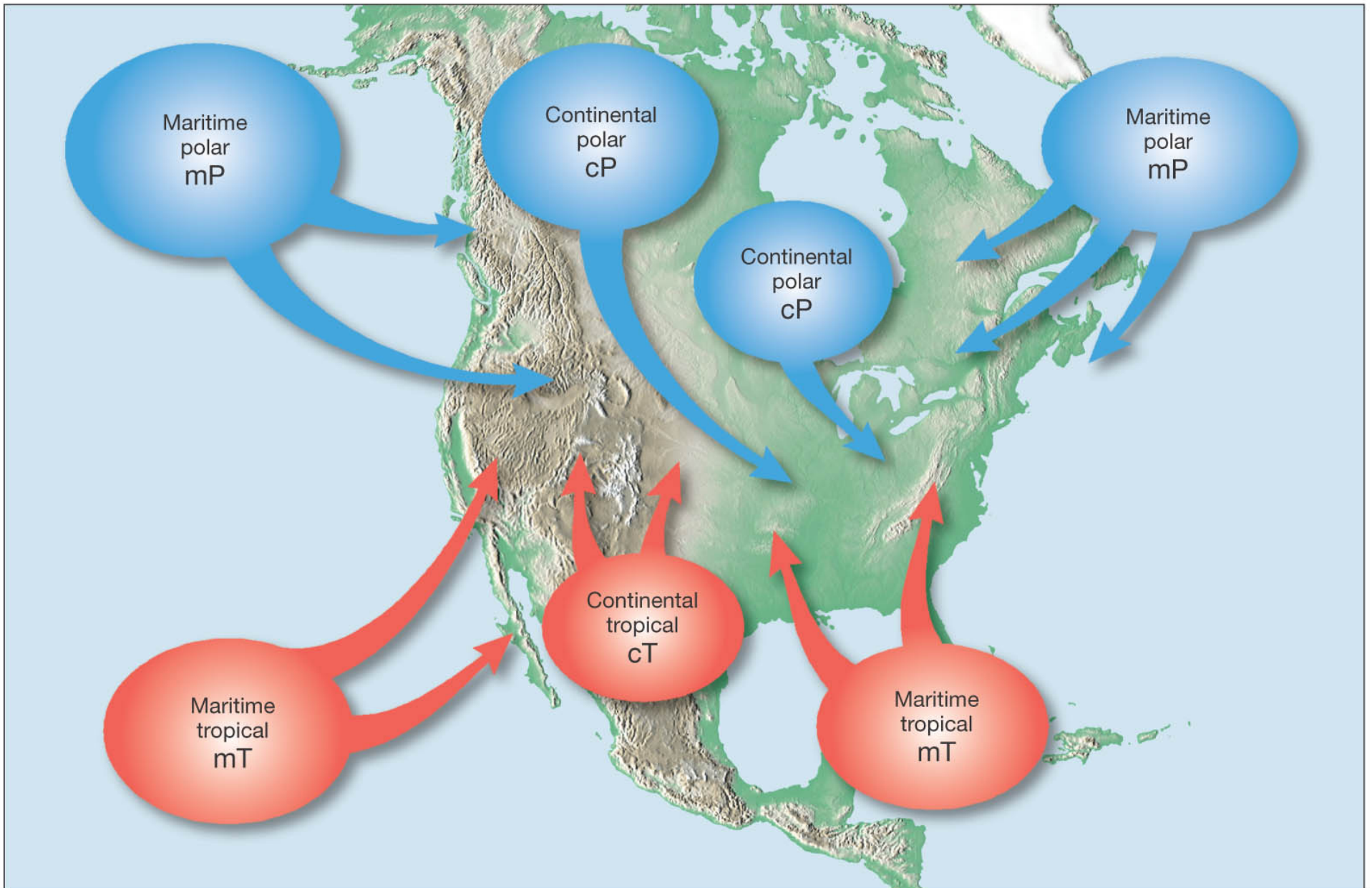
Air Masses

Extensive body of air (100's of miles horizontal extent) with mostly uniform temperature, moisture content, and atmospheric pressure

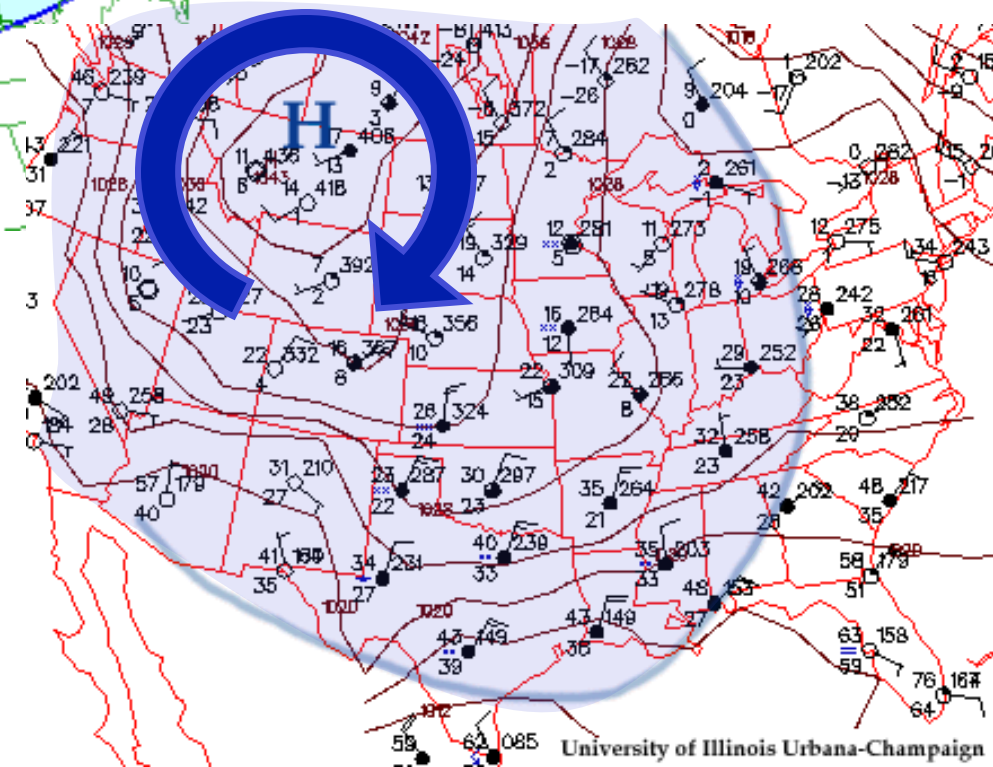
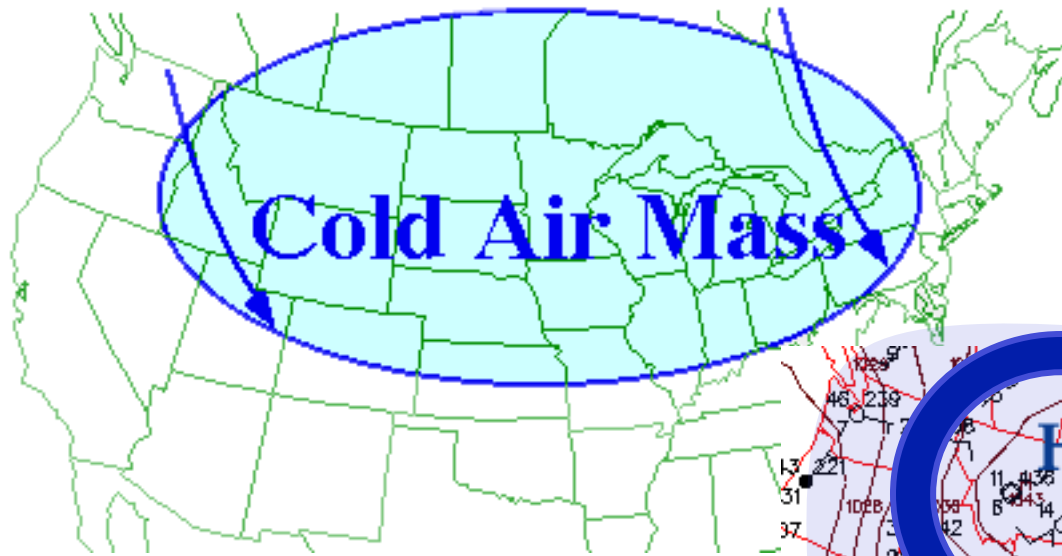
Interactions of air masses cause many kinds of large weather system.



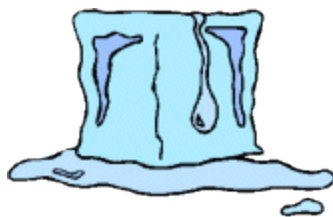


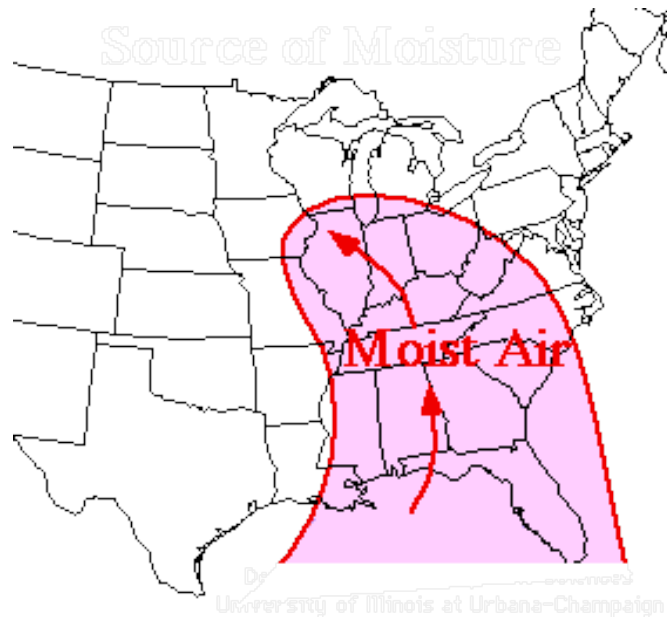


Continental Polar Air Mass



high pressure, cold, dry

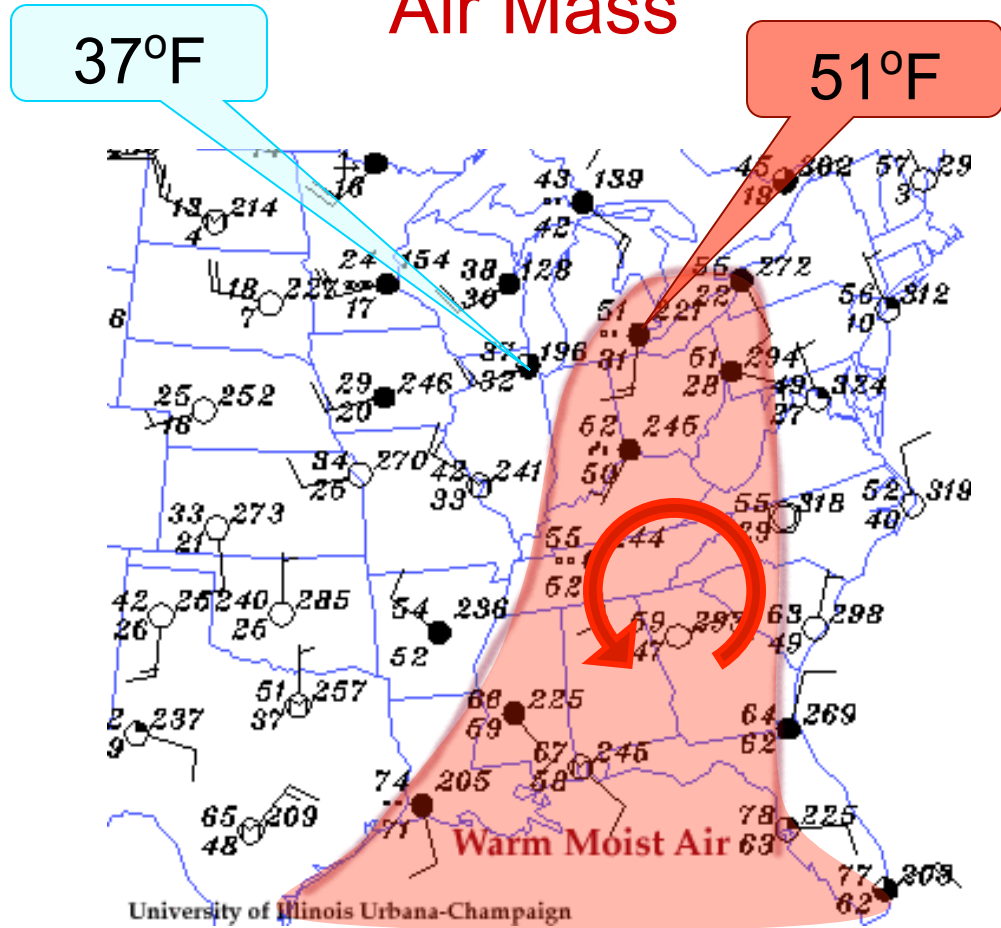




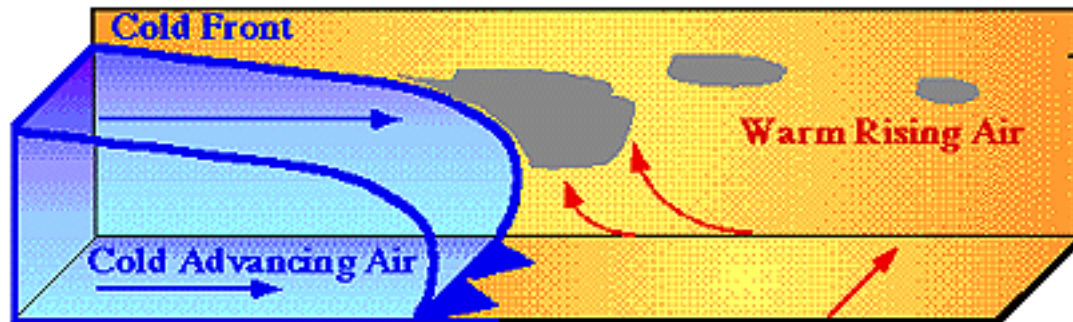
low pressure, warm, wet



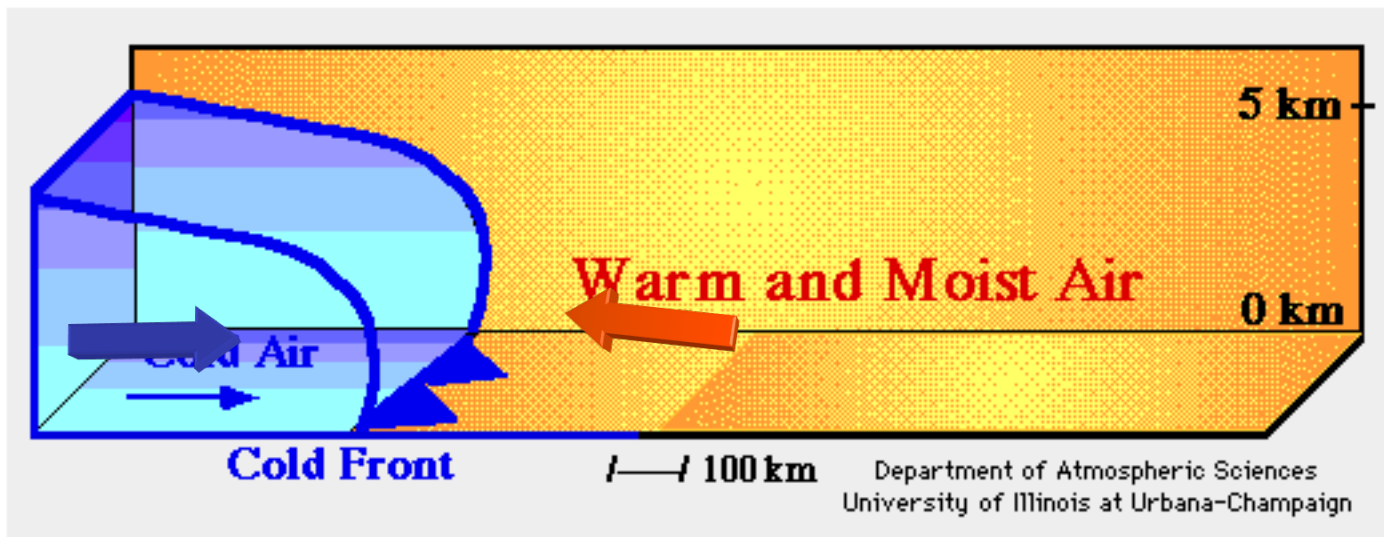
Maritime Tropical Air Mass



Front - interaction between two air masses with different characteristics (pressure, temperature, moisture content)

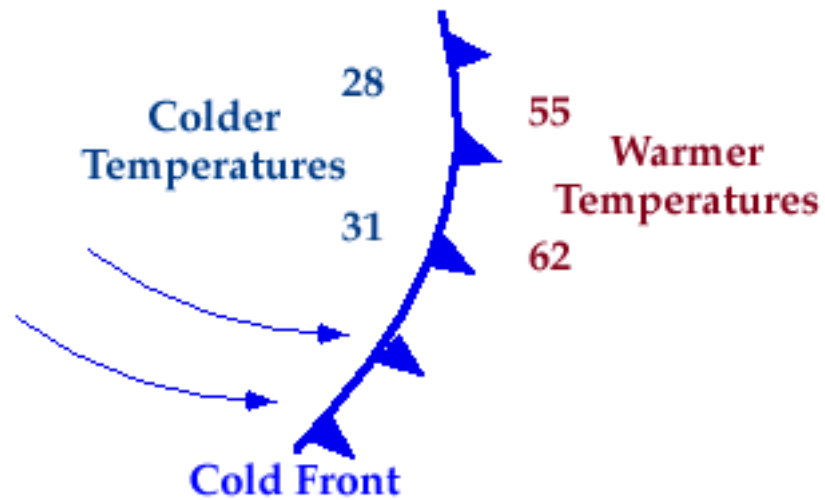


Cold Front



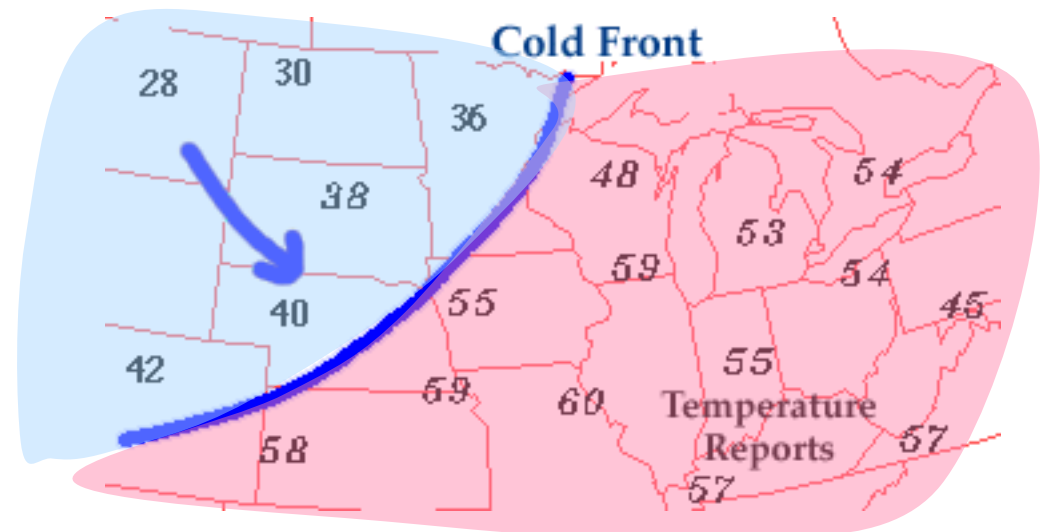
Cold air mass moves through warm air mass

Cold Front

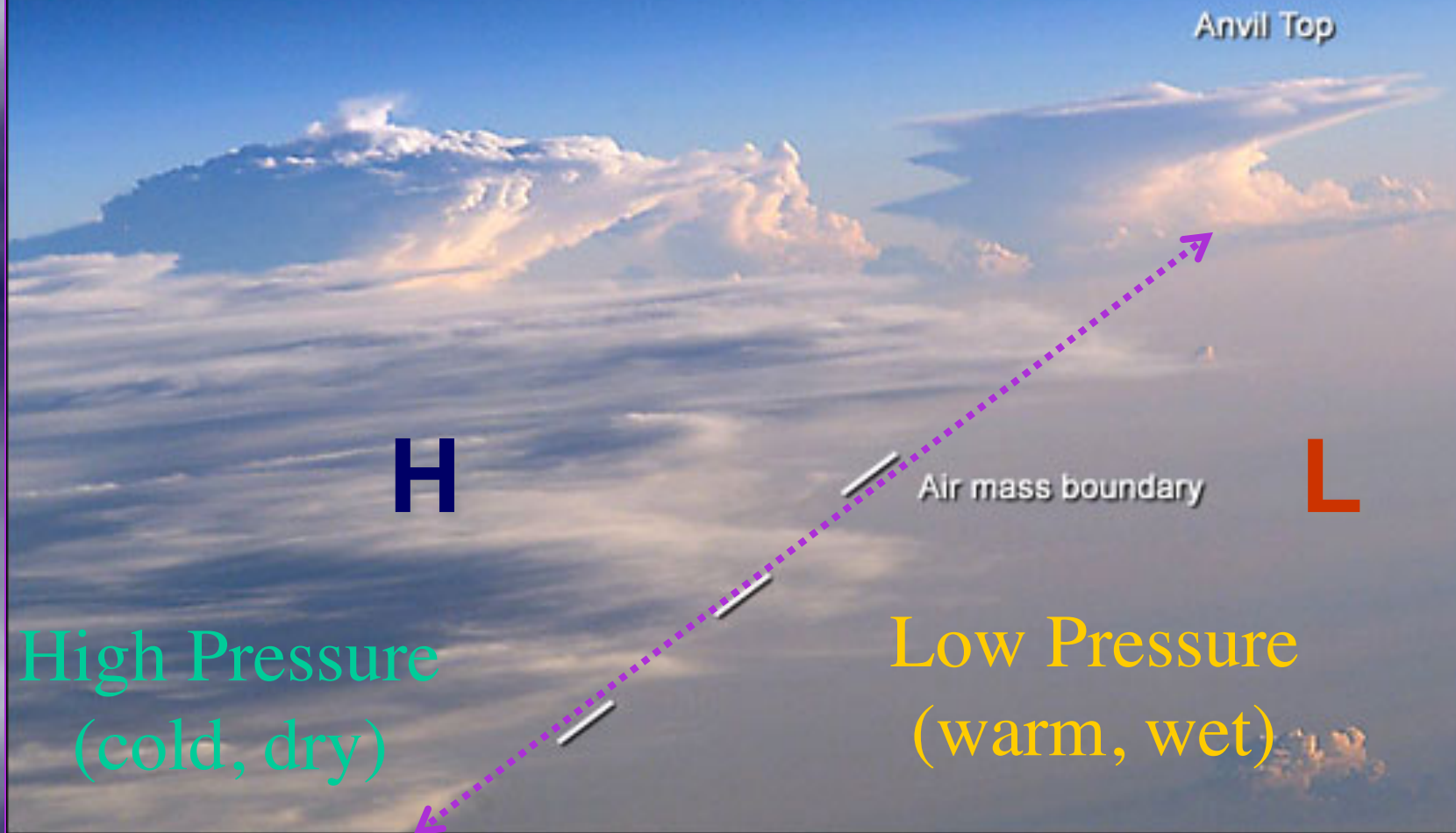


Moving front in which cold air mass replaces warm air mass

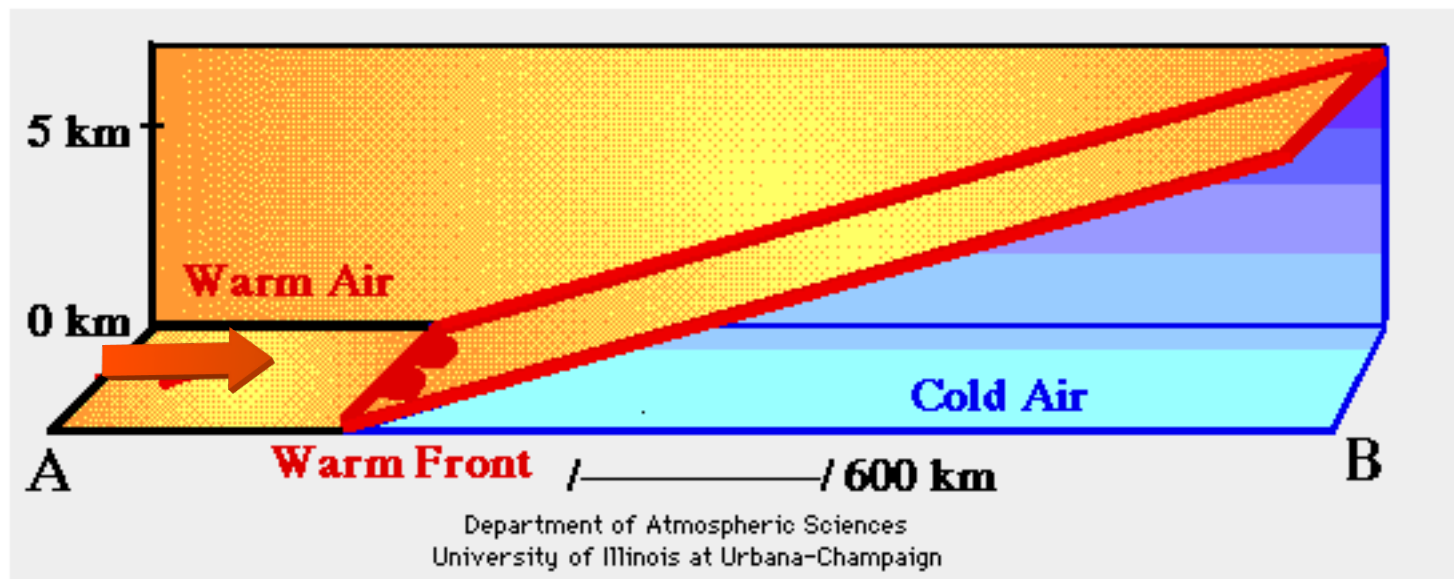
Sharp drop in temperature leads to precipitation and storms as front passes, followed by clear weather.



Cold Front



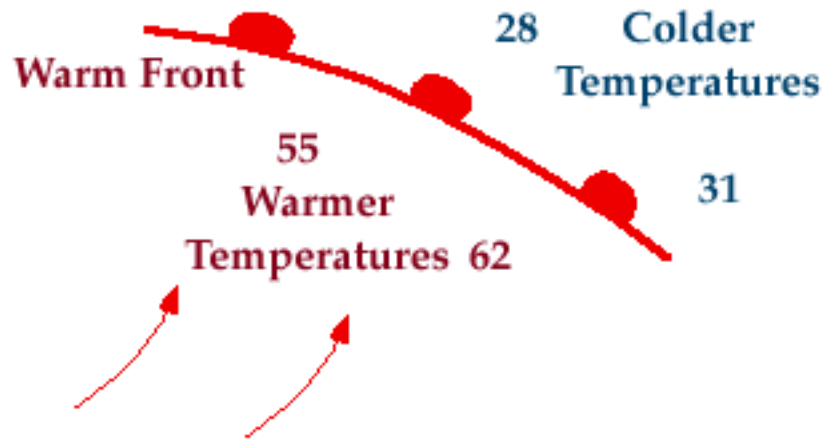
Warm Front



Warm air mass moves through cold air mass

Warm Front

Moving front in which warm air mass replaces cold air mass



Mixture leads to precipitation before and during frontal pass, followed by hazy and sometimes drizzly weather.

